

ERA-Interim bias correction of satellite radiance brightness temperatures

Algorithm, performance, and limitations

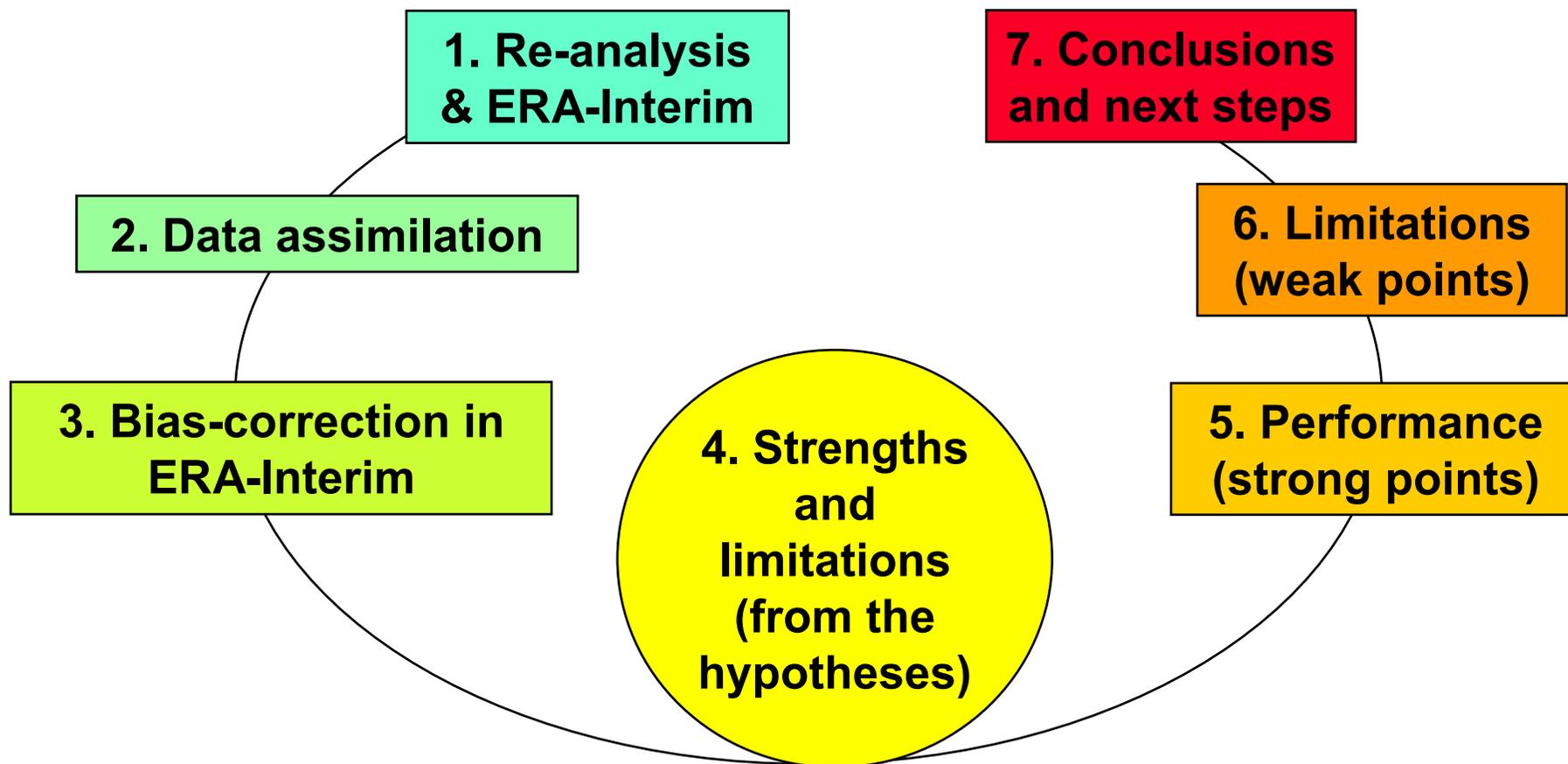


P. Poli and D. P. Dee

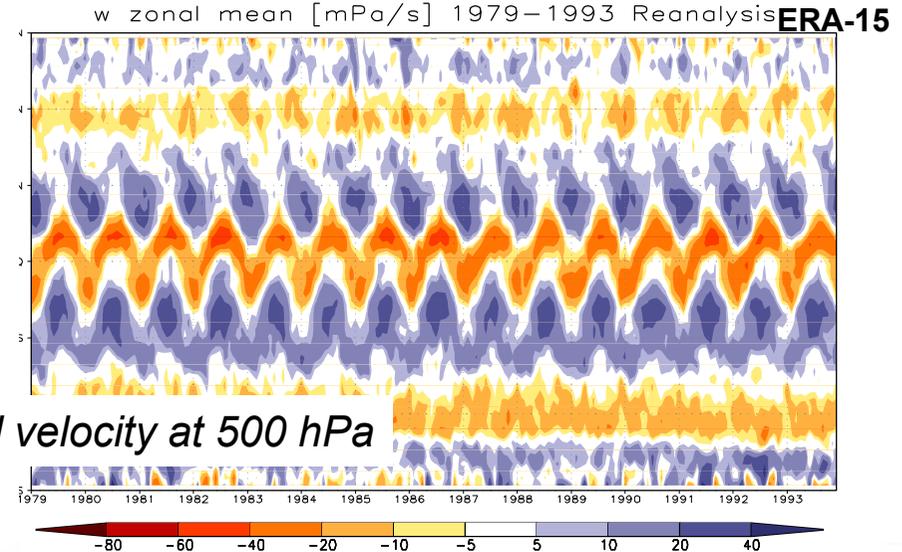
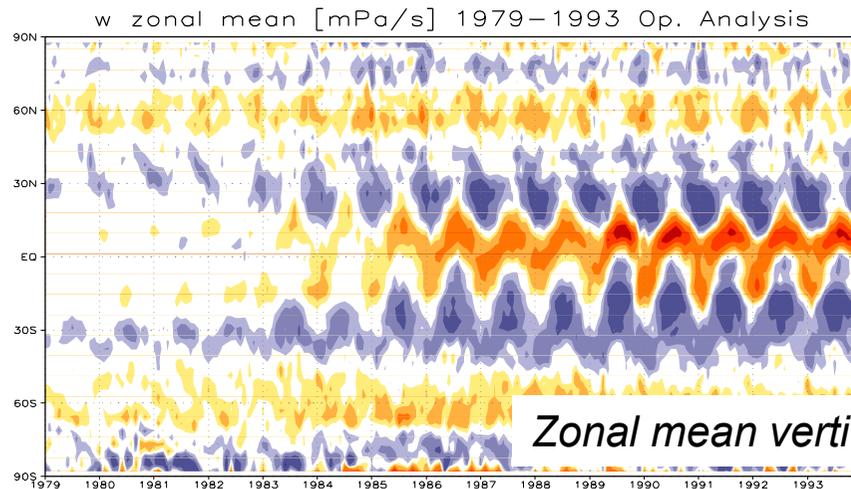
Contact: paul.poli -at- ecmwf.int

Thanks to NOAA for enabling participation in this workshop

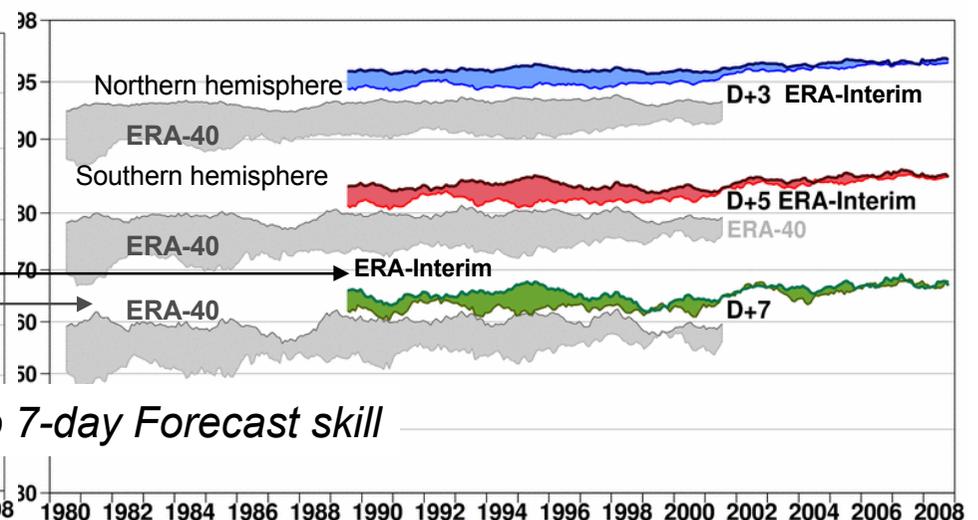
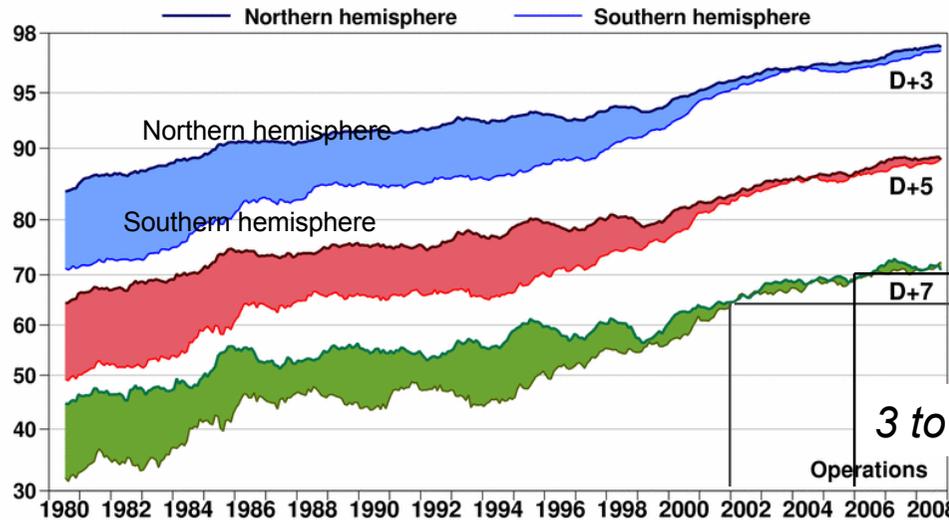
Outline (map view)



Operational NWP products vs. reanalysis products

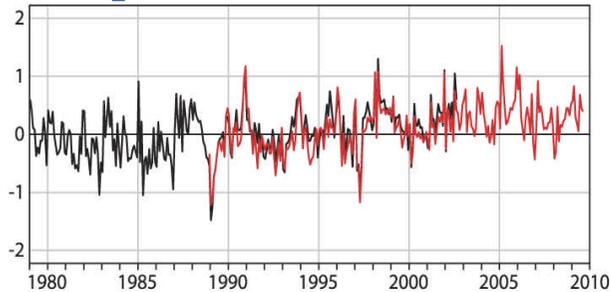


Anomaly correlation of 500hPa height forecasts



Reanalysis quality is more uniform in time than that of operational products

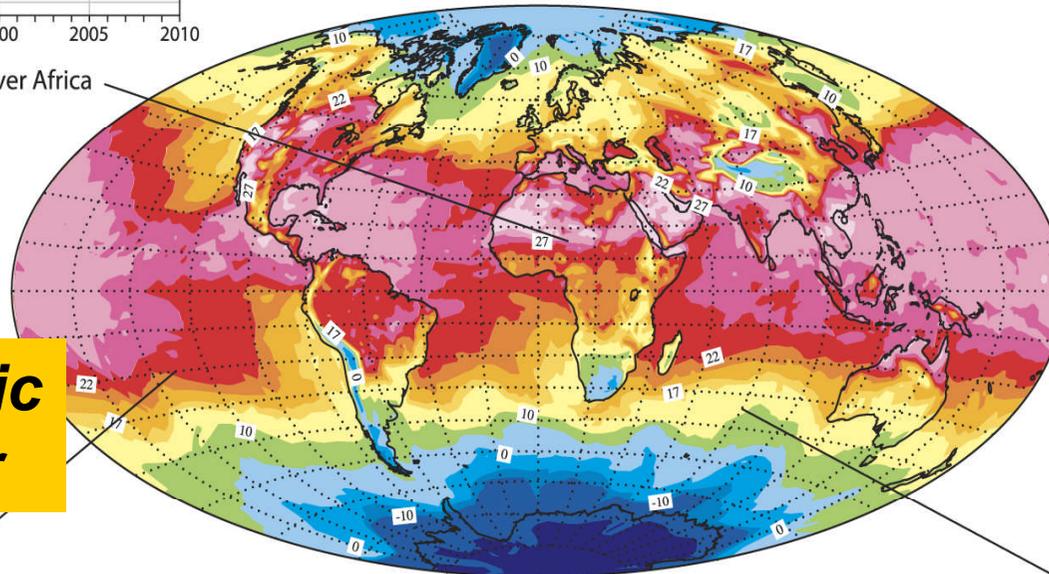
Re-analyses attempt to create a consistent picture of the Earth atmosphere



2-metre temperature anomaly (°C) over Africa

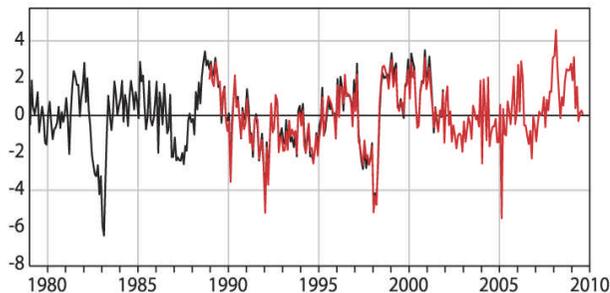
(Time)

ERA-Interim 2-metre temperature (°C)
15 August 2003 03 UTC

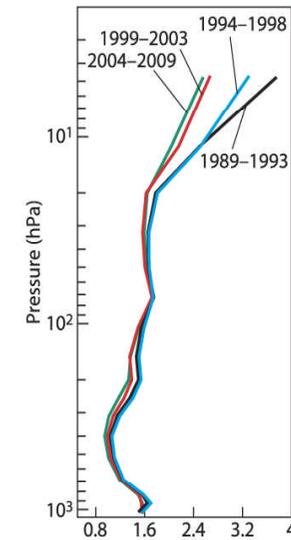


Atmospheric parameter

Southern Oscillation Index (hPa)



Horizontal



Standard deviation of differences between ERA-Interim and radiosondes temperature (°C) in the southern hemisphere

Vertical

Latest ECWMF re-analysis: ERA-Interim

- **January 1989-present** **Continues in near-real-time**
Now updated monthly
- **Horizontal resolution** T255 (~**80 km** or 0.7 deg)
- **Vertical resolution** L60 (**top 0.1 hPa** ~65 km altit.)
- **Temporal resolution:** 6 hr for upper-air fields
3 hr for surface fields
Model time-step: 30 minutes
- **Access:**
 - ECMWF member states: full archive access (MARS)
 - US education and research users: access via UCAR
<http://dss.ucar.edu/datasets/ds627.0/>
 - All users: web access via ECMWF Data Server
http://data-portal.ecmwf.int/data/d/interim_daily/

ERA-Interim Data Server

Information: <http://www.ecmwf.int/research/era>

Products: http://data-portal.ecmwf.int/data/d/interim_daily/

Data available in NetCDF or GRIB

The screenshot shows the ERA-Interim Data Server interface. The main content area is titled "ERA Interim, Daily Fields". It features a "Select date" section with a date range selector (1989-01-01 to 2006-12-31) and a "Select a list of month:" section with a grid of checkboxes for each month from 1989 to 2006. The "Select parameters" section displays a grid of checkboxes for various meteorological variables across different years and months. The variables listed include Cloud cover, Cloud ice water content, Cloud liquid water content, Divergence, Geopotential, Ozone mass mixing ratio, Potential vorticity, Relative humidity, Specific humidity, Temperature, U velocity, V velocity, Vertical velocity, and Vorticity (relative). The grid shows that many variables are selected for the years 1999, 2000, 2001, 2002, 2003, 2004, 2005, and 2006, and for the months of January, February, March, April, May, June, July, August, September, October, November, and December.

General characteristics of ECMWF re-analyses (so far)

- Aim to construct the *best* estimate of the state of the atmosphere at any given time, any given location
- Use the techniques of Numerical Weather Prediction (NWP) as employed at ECMWF
- N times daily (N=2 in ERA-Interim), run an assimilation (now **4DVAR**) to determine the most likely state of the atmosphere at a given time, so-called **analysis**
- This state is then propagated in time until the next analysis time, at which point a new data assimilation procedure is carried out
- The **consistency across variables, in space, and in time (during 12-hour intervals)** is thus ensured by the atmospheric model and its error characteristics as specified in the assimilation

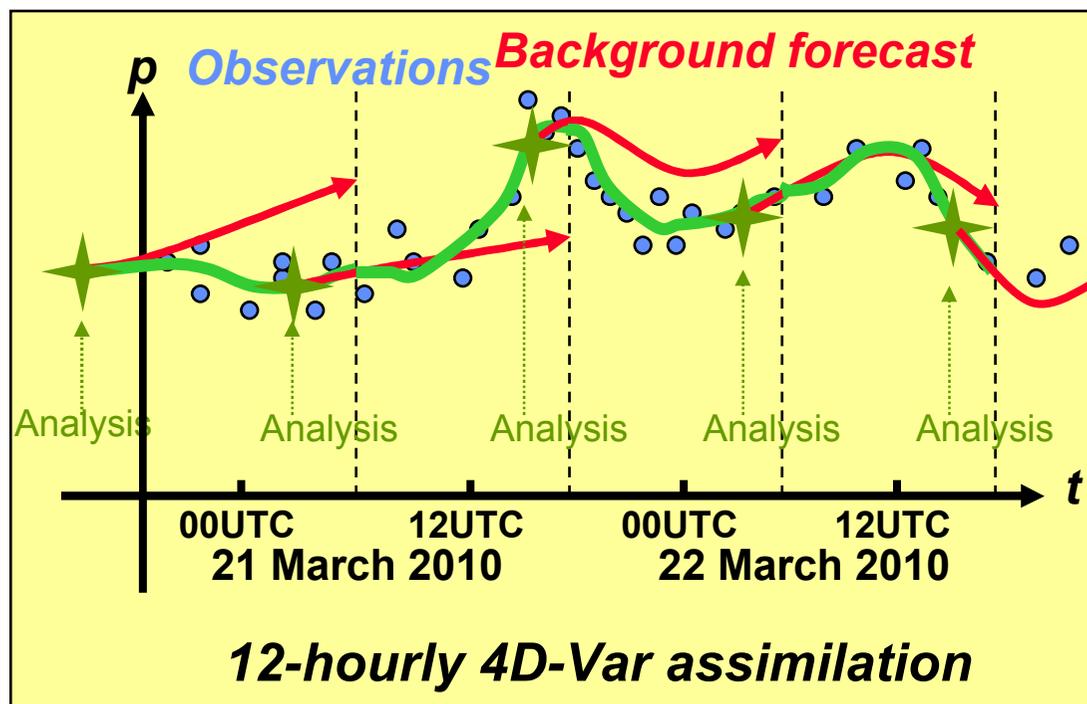
Brief primer on data assimilation

- Combine information from

- Observations
- Background forecast (propagates the information extracted from prior observations)
- Error statistical models
- Relationships to build-in dynamical and physical consistency between various meteorological parameters

- To produce the “most probable” estimate of the atmospheric state

- And some estimate of uncertainty



There's no free lunch... Hypotheses are required to get the smoothing illustrated above...

Determination of the analysis state

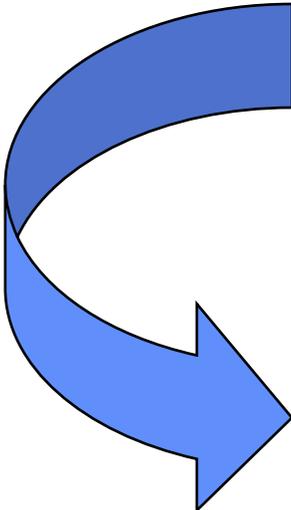
$$\text{Minimise } J(\mathbf{x}) = \underbrace{(\mathbf{x}_b - \mathbf{x})^T \mathbf{B}^{-1} (\mathbf{x}_b - \mathbf{x})}_{\text{background constraint } (\mathbf{J}_b)} + \underbrace{[\mathbf{y} - \mathbf{h}(\mathbf{x})]^T \mathbf{R}^{-1} [\mathbf{y} - \mathbf{h}(\mathbf{x})]}_{\text{observational constraint } (\mathbf{J}_o)}$$

- The input x_b represents past information propagated by the atmospheric forecast model (the **background**)
- The input $[y - h(x_b)]$ represents the new information (observations) entering the system (the **background departures**)
- The function $h(x)$ represents a model for simulating observations; this includes the forecast model in 4DVAR (the **observation operator**)
- Minimising the cost function $J(x)$ produces an adjustment to the model background based on all used observations (the **analysis**)
- **Hypotheses typically used to solve this** (and/or buried in this approach):
The background and the observation errors are independent, Gaussian with zero means (**unbiased**), and properly specified in B and R; the model errors are negligible within the analysis window.

Variational bias correction

First proposed and implemented by Derber and Wu, 1998

- Solve for analysis and bias parameters at the same time


$$\text{Minimise } J(\mathbf{x}) = (\mathbf{x}_b - \mathbf{x})^T \mathbf{B}^{-1} (\mathbf{x}_b - \mathbf{x}) + [\mathbf{y} - \mathbf{h}(\mathbf{x})]^T \mathbf{R}^{-1} [\mathbf{y} - \mathbf{h}(\mathbf{x})]$$

The **bias parameters**: β

The **bias model**: $b(\mathbf{x}, \beta)$: Typically a linear combination of bias parameters with robust predictors to characterize air mass or observation geometry

$$\text{Minimise } J(\mathbf{z}) = (\mathbf{z}_b - \mathbf{z})^T \mathbf{B}_z^{-1} (\mathbf{z}_b - \mathbf{z}) + [\mathbf{y} - \tilde{\mathbf{h}}(\mathbf{z})]^T \mathbf{R}^{-1} [\mathbf{y} - \tilde{\mathbf{h}}(\mathbf{z})]$$

$$\mathbf{z}^T = [\mathbf{x}^T \beta^T] \quad \tilde{\mathbf{h}}(\mathbf{z}) = \mathbf{h}(\mathbf{x}) + \mathbf{b}(\mathbf{x}, \beta)$$

- The aim is to correct for observation *and* observation operator (radiative transfer) error bias – altogether
- Assuming that these biases are constant during the duration of the analysis window

Bias predictors in ERA-Interim for radiances

Sensor	Bias predictors											
	offset	Δz 1000 -300mb	Δz 200 -50 mb	Δz 10 -1 mb	Δz 50 -5 mb	Total col wv	Skin Temp.	Sfc wind speed	Nadir view angle NVA	NVA **2	NVA **3	NVA **4
HIRS												
AIRS												
GEO IMG												
SSU	Except channel 3											
MSU												
AMSU-A	Except channel 14											
AMSU-B												
MHS												
SSM/I												
SSM/I-S												
AMSR-E												

Central node of this talk: Strengths and limitations of the variational bias correction

● Strengths:

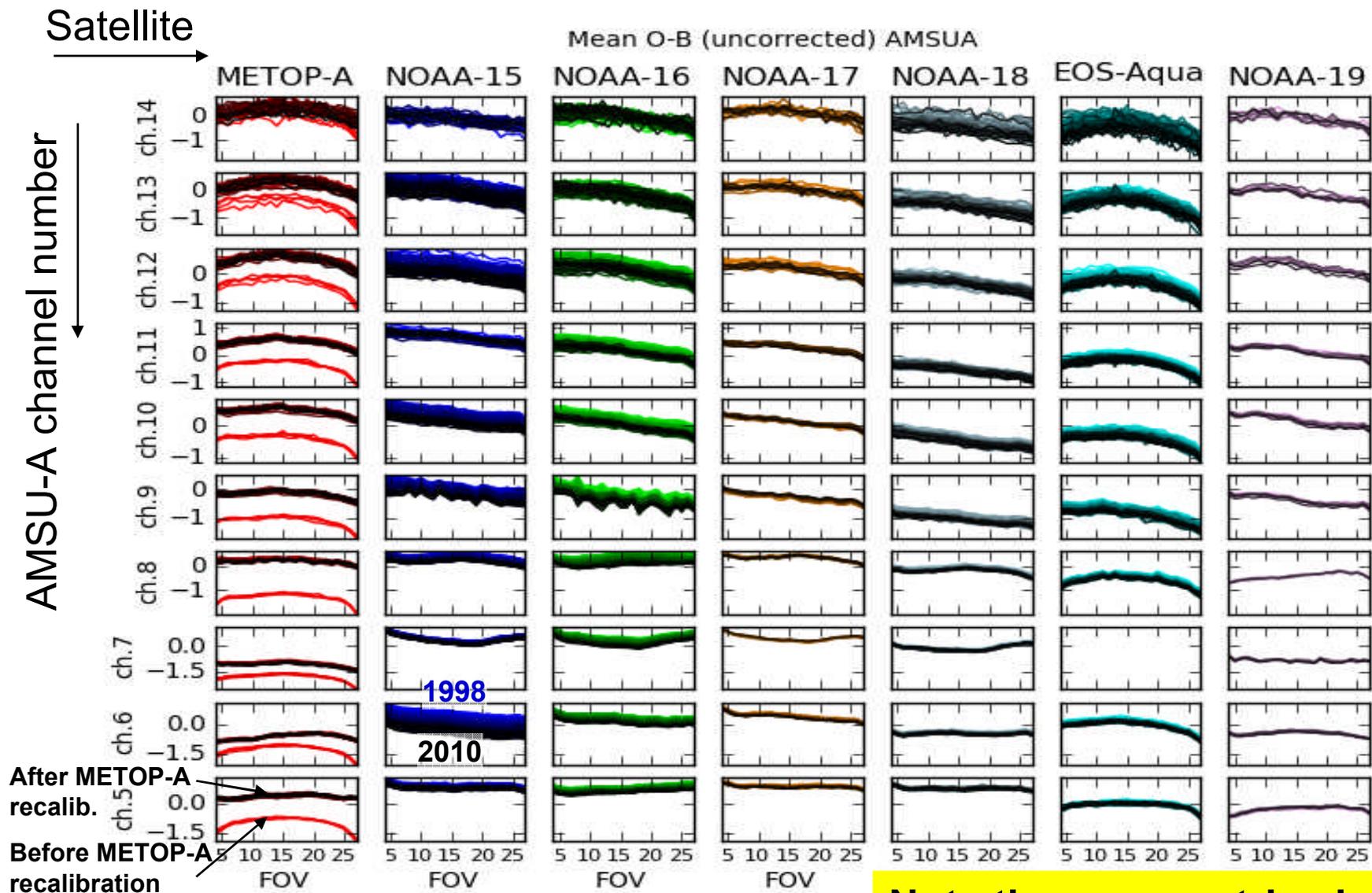
- Benefit from all the other observations to determine the most likely bias estimate
- “Quick” updates (12 hours) are possible
- Automatic procedure
- Do not assume that observations present consistent bias characteristics over the life-time of each instrument

● Limitations:

- When/where there are few observations with independent biases, or
- When/where the only observations available present similar biases as compared to the model,
- Then the model bias can contaminate the bias estimate

Example 1:

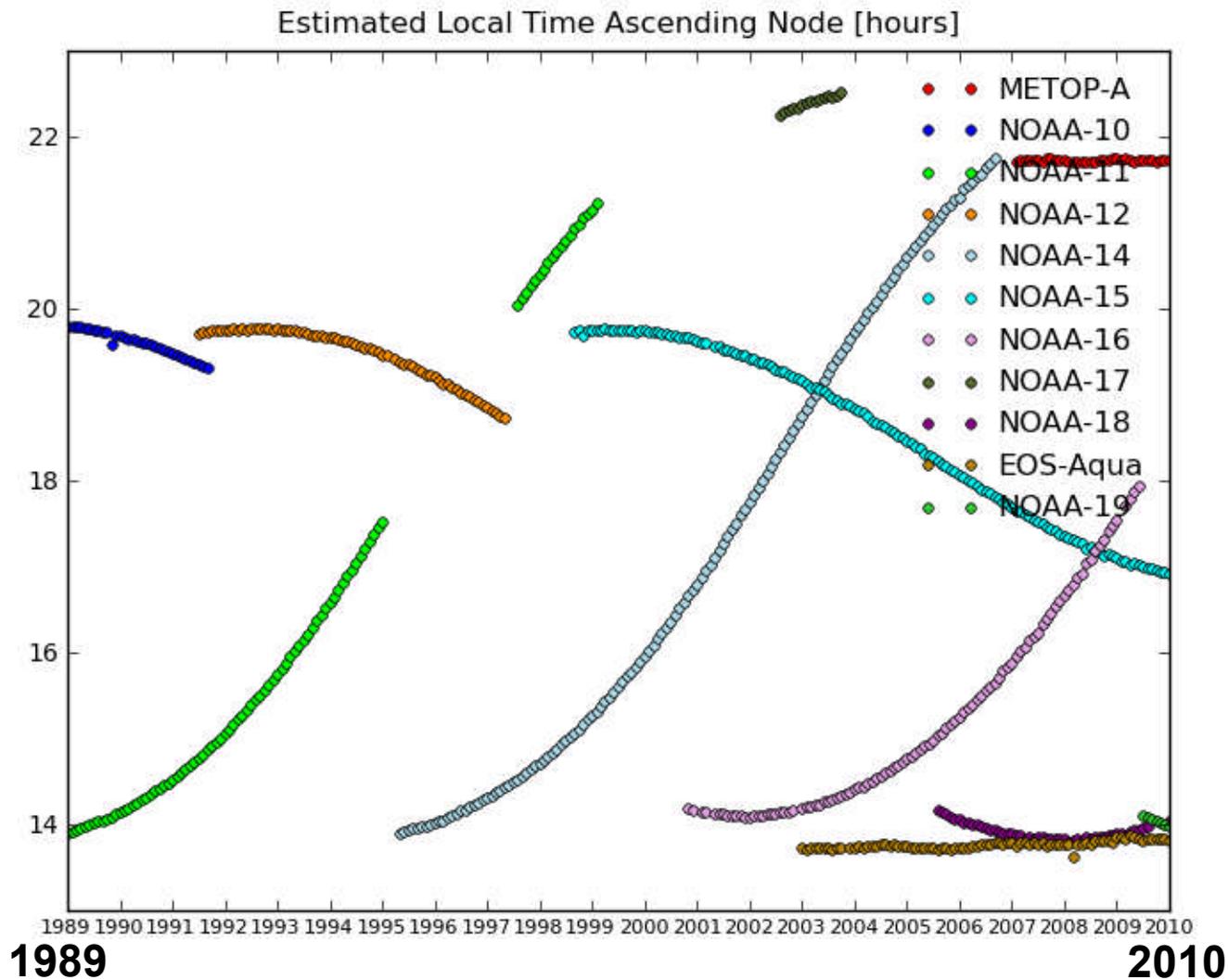
AMSU-A Scan-angle bias



Note the asymmetric shapes

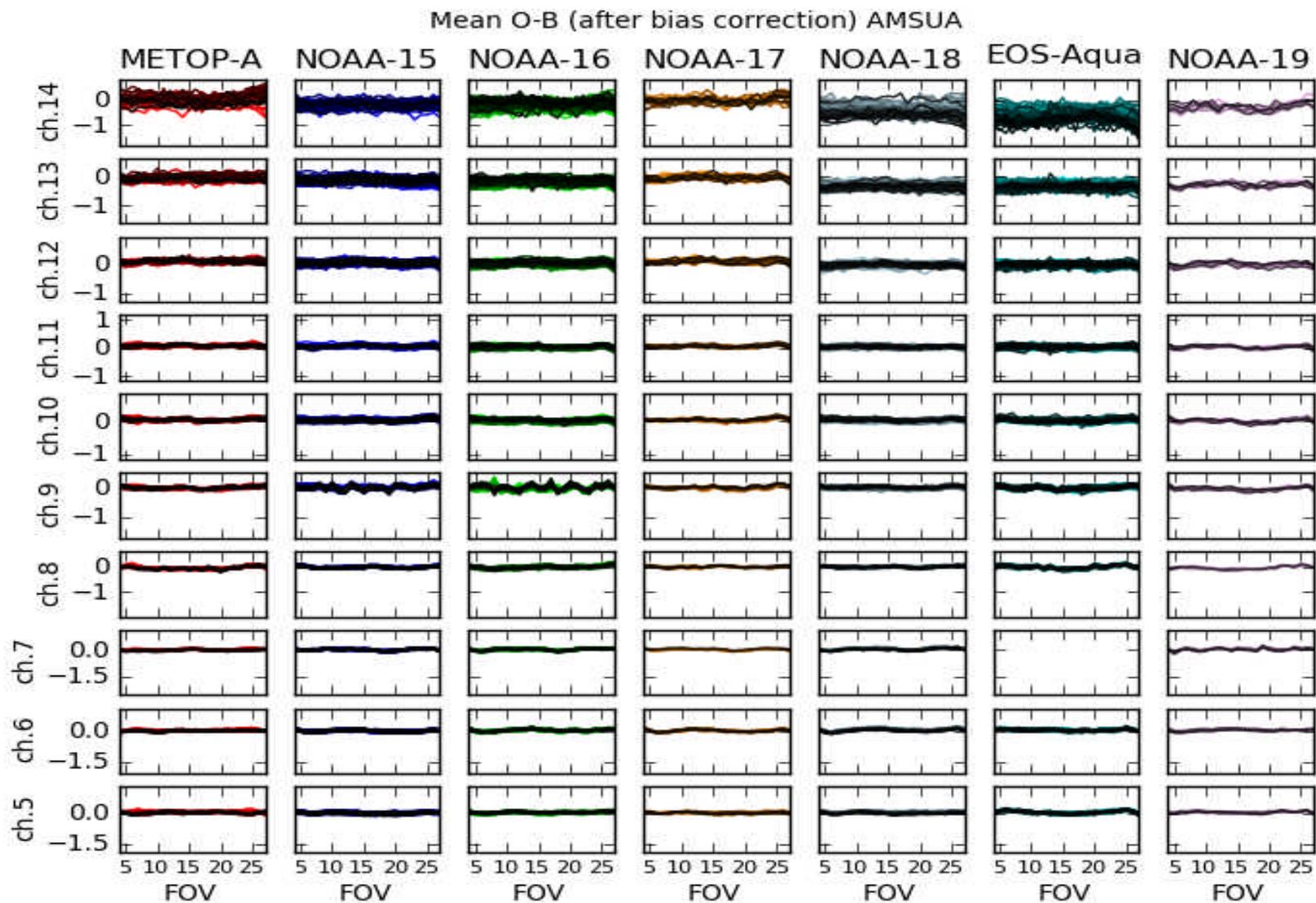
Latitudes 20S-20N, ocean only

Satellite orbital drift, as estimated from the data at 1-month intervals



Example 1 (cont):

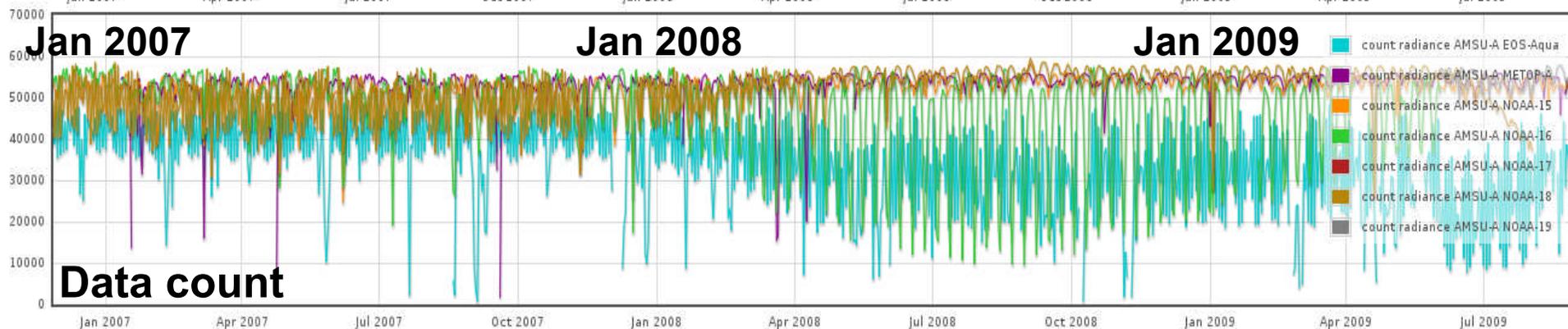
AMSU-A biases, after variational bias corr.



Example 2:

AMSU-A METOP-A recalibration after launch

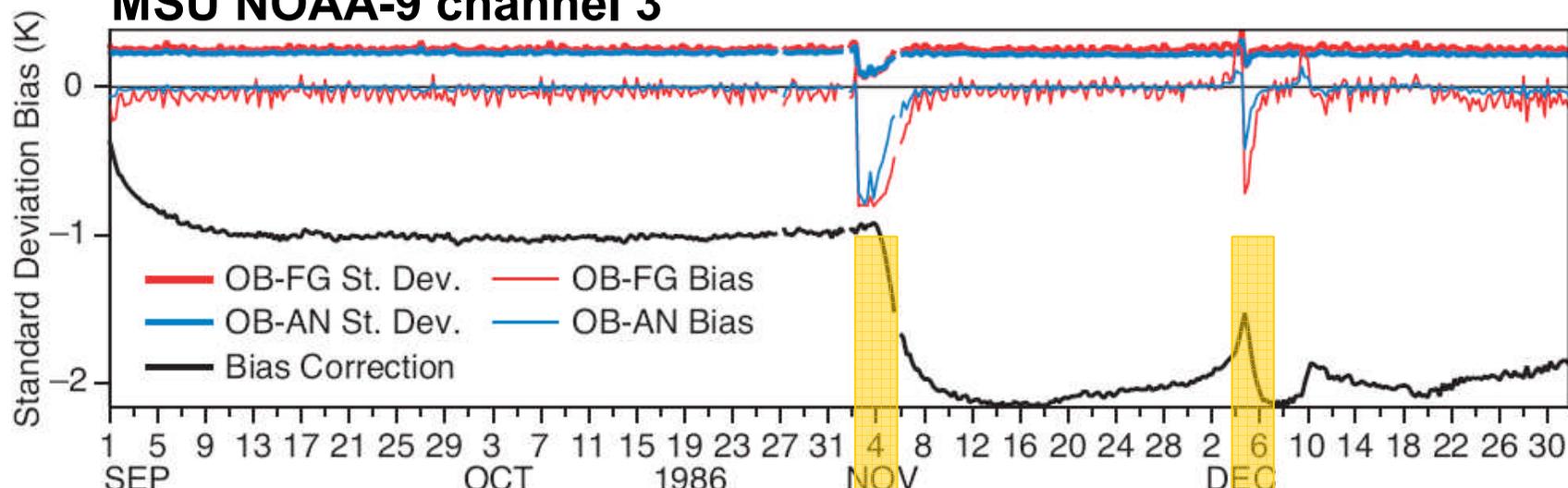
Mean global bias correction (K) for channel 9



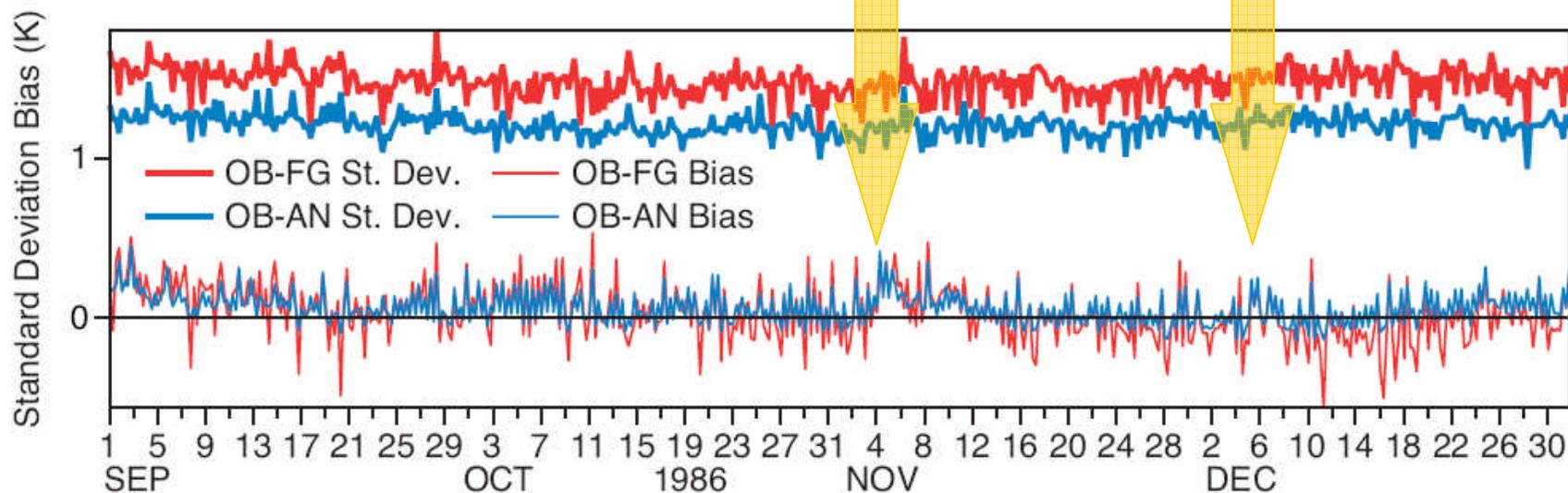
Example 3:

Magnetic storms

MSU NOAA-9 channel 3

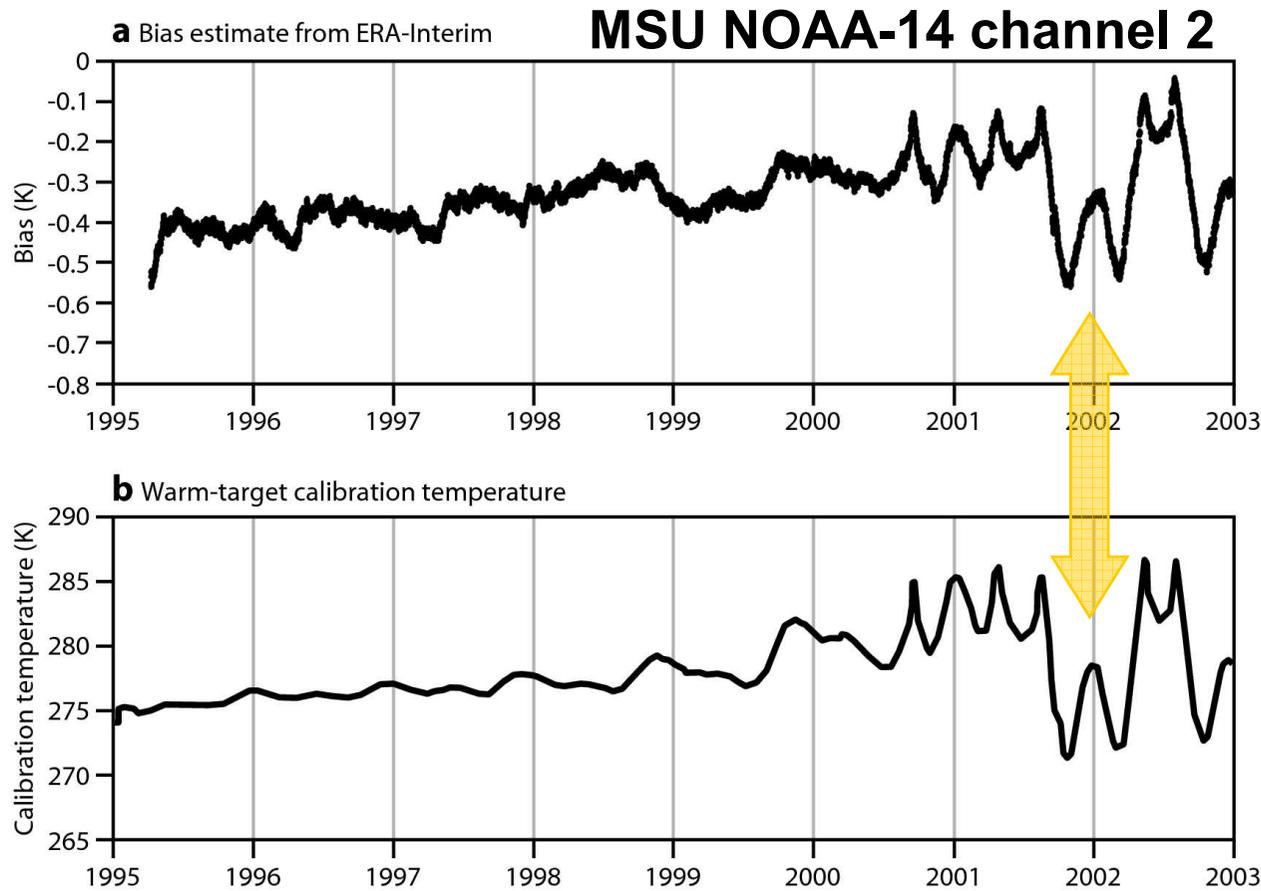


200 hPa temperature departures from radiosonde observations



Example 4:

Reference blackbody calibration fluctuations



Variational bias estimates for NOAA-14

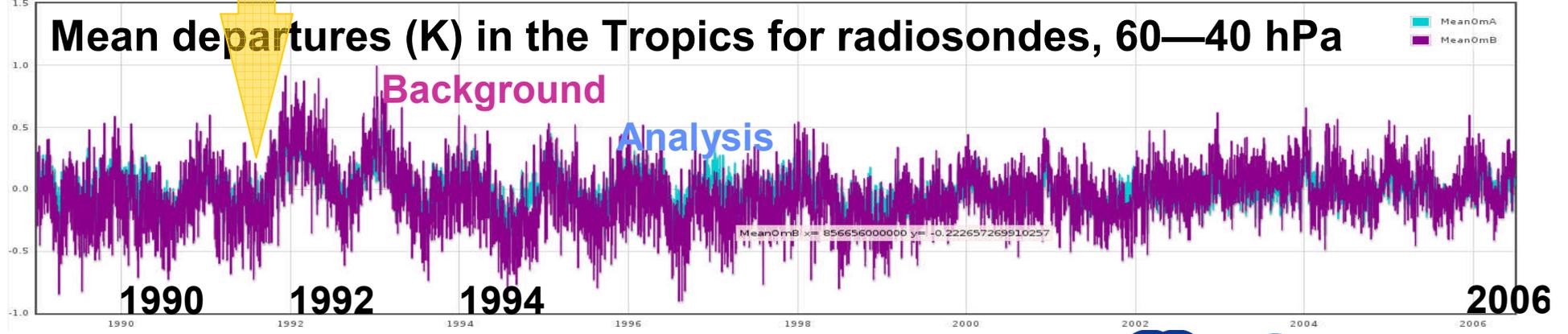
Actual warm-target temperatures on board NOAA-14 (Grody *et al.* 2004)

Dee and Uppala, 2009

Example 5 :

Mt Pinatubo eruption

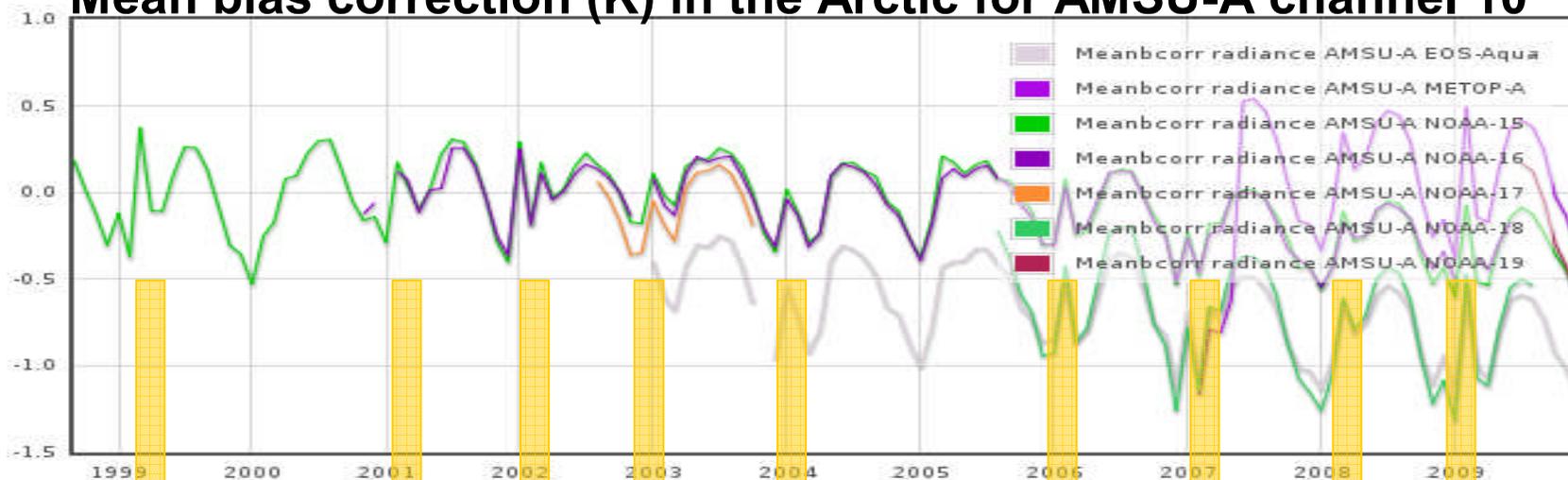
Mean obs-background departures (K) in the Tropics for MSU channel 4



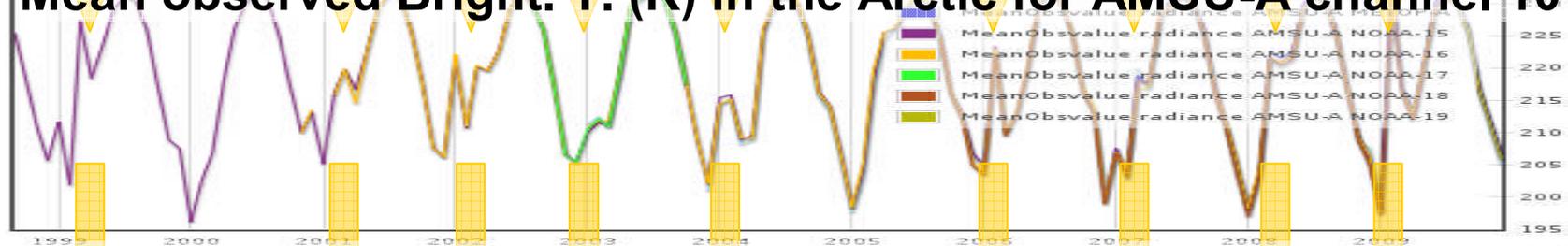
Example 6:

Sudden stratospheric warming events

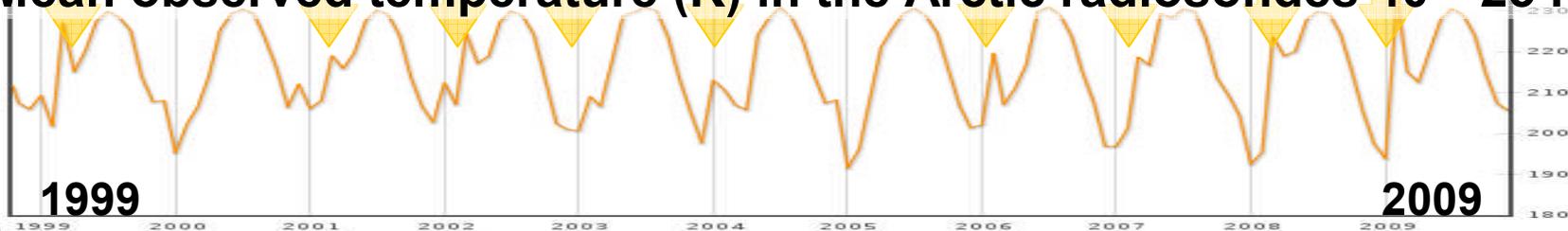
Mean bias correction (K) in the Arctic for AMSU-A channel 10



Mean observed Bright. T. (K) in the Arctic for AMSU-A channel 10



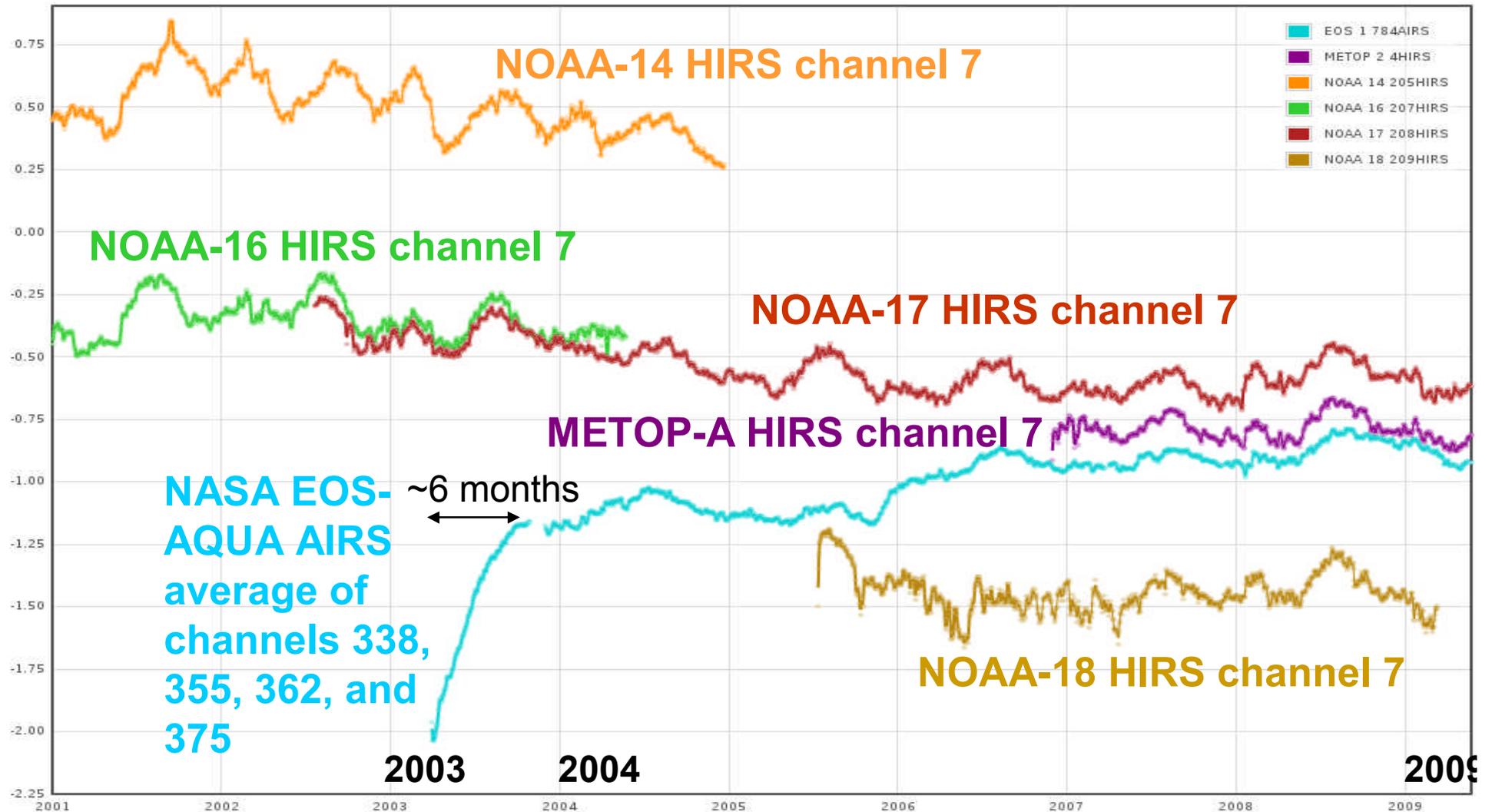
Mean observed temperature (K) in the Arctic radiosondes 40—25 hPa



Example 7:

Slow spin-up for AIRS

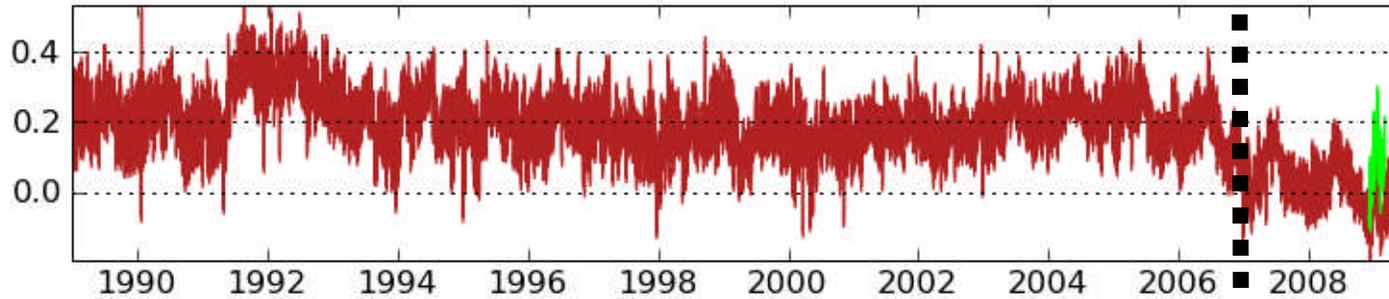
Mean global bias correction (K)



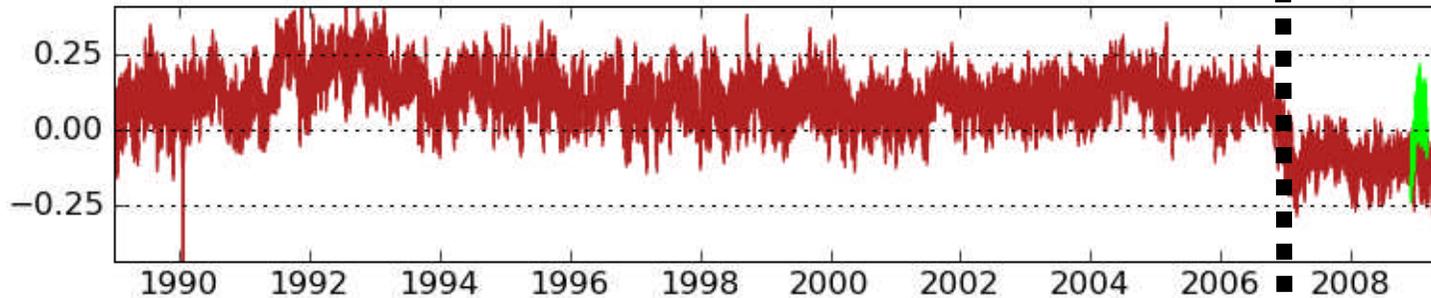
Example 8:

Sensitivity to changes in the “anchoring” obs. system: GPSRO

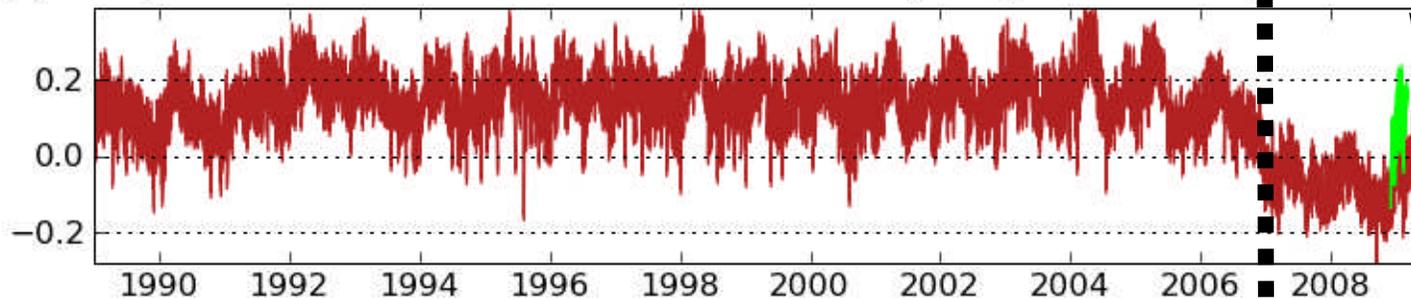
(a) Temper. diff. NH land RS minus ERA-Interim (in K), Pressure layer 60-40hPa



(b) Temper. diff. NH land RS minus ERA-Interim (in K), Pressure layer 85-60hPa



(c) Temper. diff. NH land RS minus ERA-Interim (in K), Pressure layer 125-85hPa

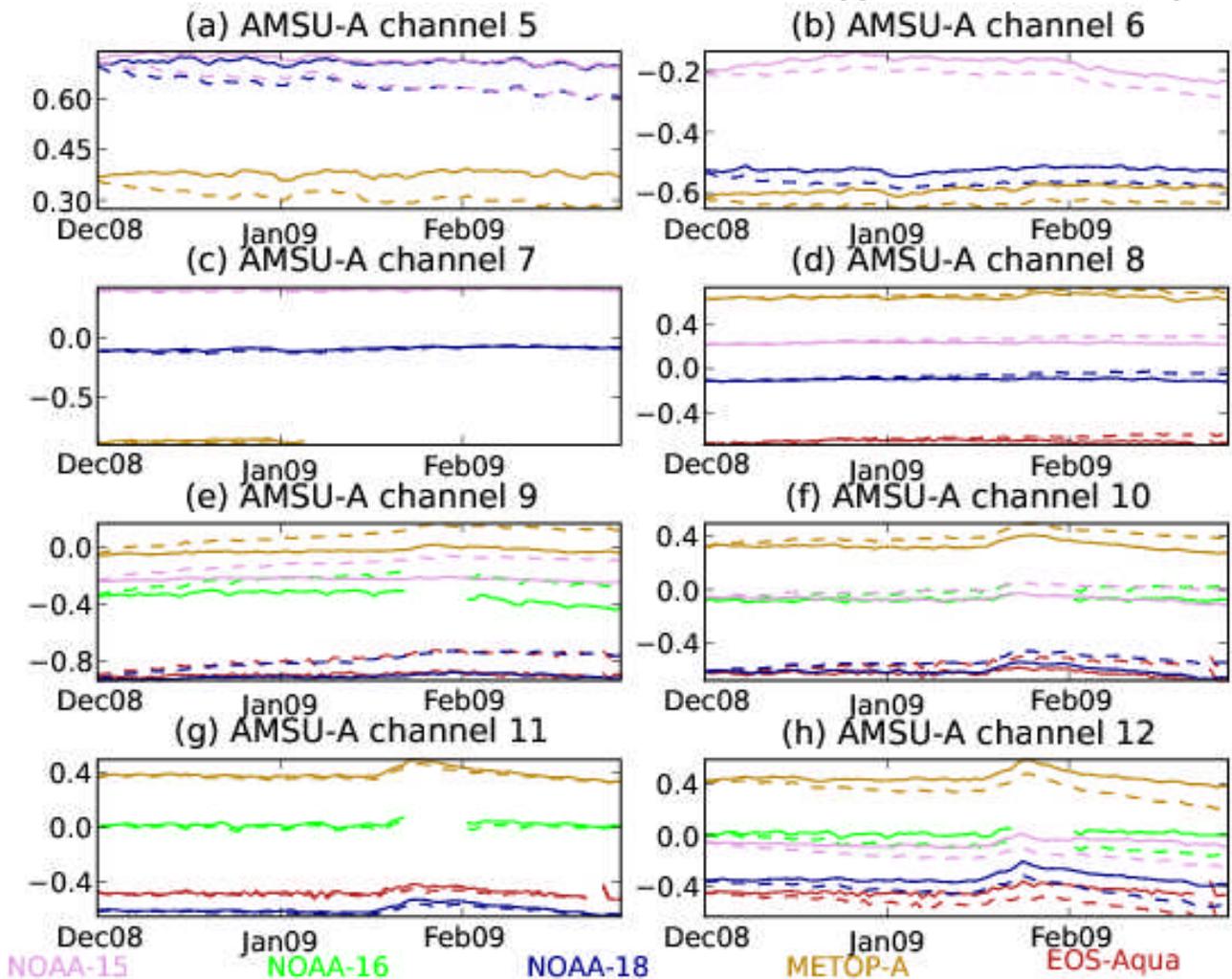


Observing System Experiment, in which GPSRO data are *not* assimilated

Introduction of GPSRO COSMIC

Example 8 (cont.): Sensitivity to changes in the “anchoring” obs. system: GPSRO

Bias correction estimates (global means)



Observing System Experiment, in which GPSRO data are *not* assimilated

— ERA-Interim - - - - ERA-Interim *without* GPSRO assimil.

Summary: Variational bias correction

(1) Weak points

- **Sensitive to changes in the “reference” obs. system:**
 - Aircraft data assimilated in ERA-Interim are biased warm, and increasingly numerous; These data need to be bias corrected in the next ERA
 - Introduction of a large number of GPSRO soundings introduces a break in the time-series of temperatures at the tropopause
- **The variational bias correction cannot do everything!**
 - “Some” handling of model biases is definitely required for the next ERA; aerosols need better handling in particular
- **Spin-up can sometimes be slow:**
 - This seems to happen when a sensor probes a sub-space not observed by other instruments; the model state then takes a while to adjust

Summary: Variational bias correction (2) Strong points

- **Successfully correct for scan-angle dependent biases**
 - These cannot be explained simply by model or radiative transfer model biases
- **Manage to anchor the reanalysis system to observations which we specify as references**
 - it's not black magic anymore where the mean behavior is the “sum” of all the observations
 - Opens the possibility to run several reanalyses where various observing systems are specified as references; the spread being an indication of how well we understand the observing systems and the past climate
- **The shocks introduced by new satellite radiances, as found in ERA-40, are greatly reduced**
- **The technical complexity of satellite data assimilation has been greatly reduced**

Next steps

- The experience gained with ERA-Interim will be valuable to feed back findings and verify whether the bias estimates are consistent with other estimates
- *Biases asymmetric with respect to scan angle seem to point to phenomena still to be better corrected -- do antenna corrections need improving?*
- Bias correction is likely to stick around for a while, but we hope to run various reanalyses with different types of observing systems as references
- Look forward to assimilating recalibrated observations – to continue the improvement loop...

Initial raw satellite record

Recalibrated satellite record

Improved CDR

Lessons v3

Lessons v2

Lessons learnt from long assimilation and comparison with other sources

Iterative process based on given raw satellite data

Reanalysis

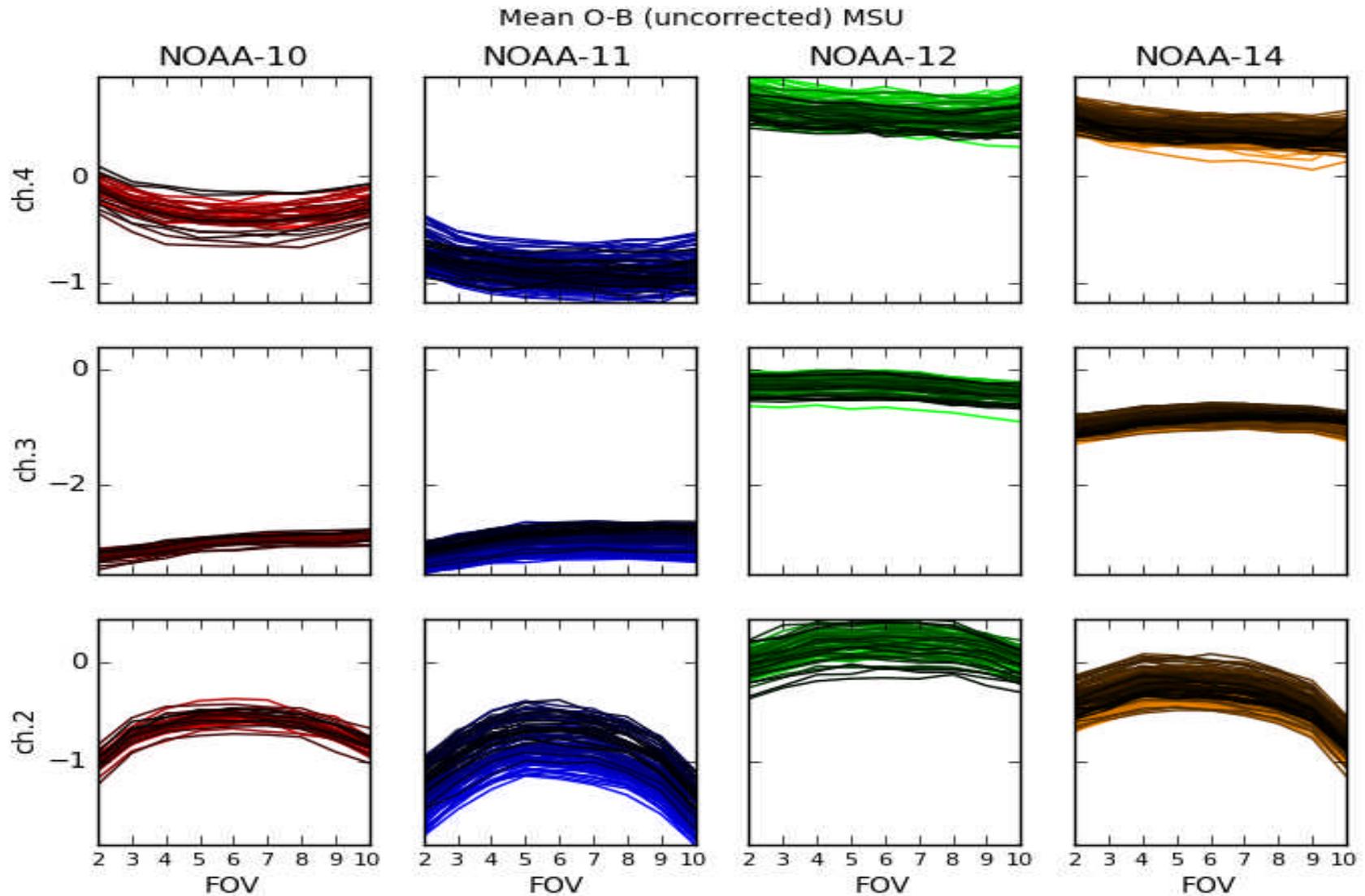
Improved reanalysis (v2)

Improved reanalyses (v3)

Climate Truth

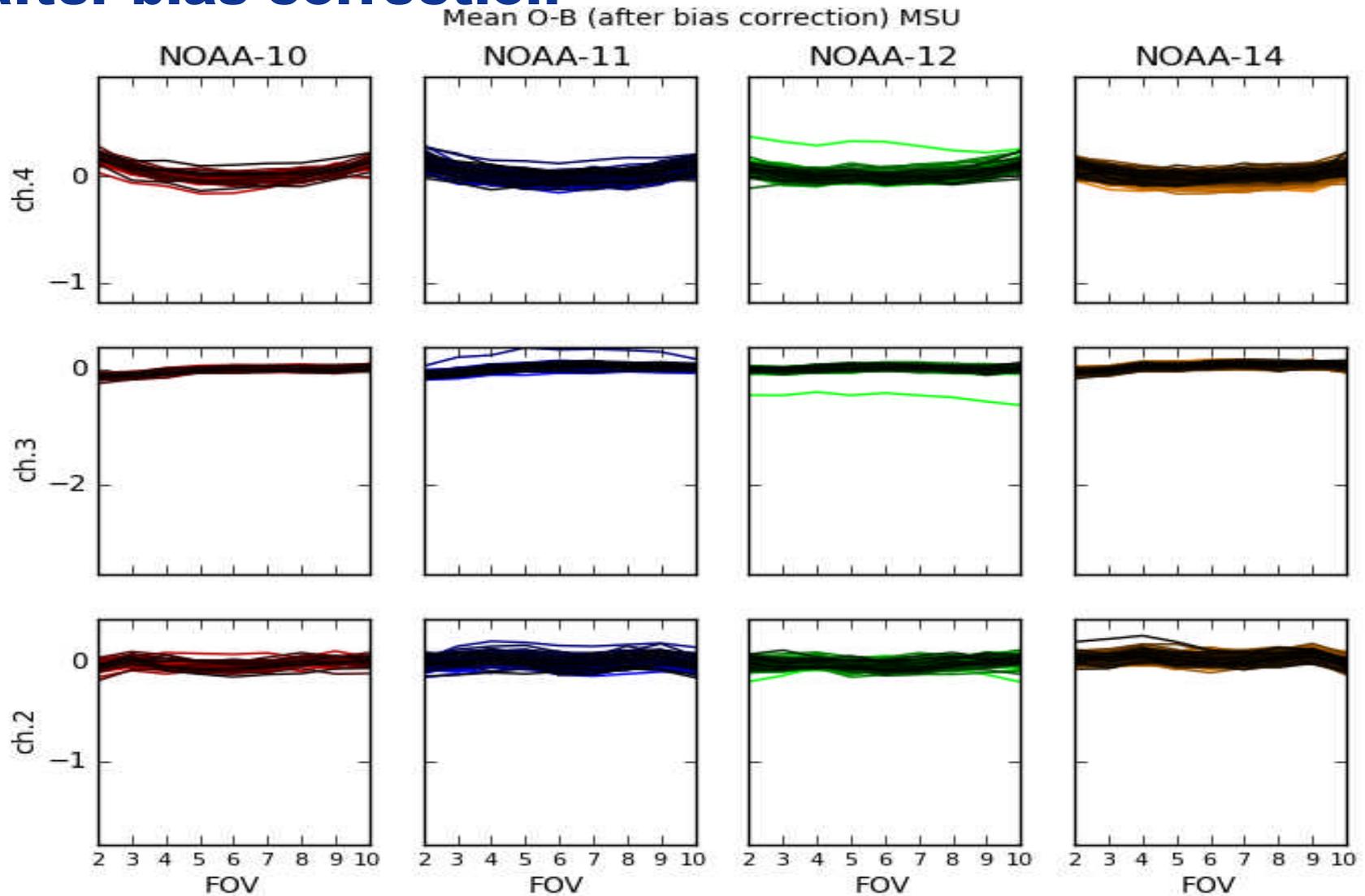
ERA-75
ERA-40
ERA-Interim

Scan-angle dependence of the MSU biases



Curves get darker as the dates increase

Scan-angle dependence of the MSU biases after bias correction



Curves get darker as the dates increase